



SURFICIAL GEOLOGY OF THE WITHERBEE 7.5-MINUTE QUADRANGLE, ESSEX COUNTY, NEW YORK

Geographic data layers from 2017 TIGER/Line shapes for transportation

Shaded relief from 2015 Clinton, Franklin and Essex County 1m, 2015 Warren, Washington and Essex County 1m, and NYS 10m DEM lidar data sets

Magnetic declination from the NOAA online Declination Calculator:

http://www.ngdc.noaa.gov/geomag-web/#declination

(http://gis.ny.gov/elevation/index.cfm)

and hydrograpghy (https://www.census.gov/cgi-bin/geo/shapefiles/index.php)

Andrew Clift and Brandon Graham 2013

1000 500 0 1000 2000 3000 4000 5000 6000 7000 8000 9000 10000

CONTOUR INTERVAL: 10 FEET

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New York State Geological Survey

2013 Digital data and cartography, A. Clift and K. Backhaus, 2013 & 2019.

Map Discussion

The Witherbee quadrangle is located along the eastern boundary between the High Peaks and Central Highlands physiographic subdivisions of the Adirondack Province in northern New York State. It contains approximately 139 km² of public and private land that is heavily forested and home to a variety of plant and animal species. Located in the east-central portion of this quadrangle is the historic town of Moriah, known for its significant role throughout the 19th and 20th centuries as a major iron mining town in New York State. The physiography of this region consists of numerous hills and northeasttrending valleys that form broad, topographic prominences throughout the landscape west of Moriah. Drainage from these valleys funnels east towards Moriah Center, which is situated at the confluence of several streams and brooks that enter into Mill Brook, and ultimately into Lake Champlain, approximately 6 km east of Moriah Center. The highest point in this quadrangle is Bald Peak at 704 m above sea level (ASL), and the lowest point is just east of Moriah Center along Mills Brook at approximately 210 m ASL.

The nature of the topography in this region is most broadly controlled by structural deformation in the bedrock that is primarily associated with the Grenville Orogeny (1,250 - 980 M.a.). Zones of bedrock weakness attributed to episodes of complex faulting, folding and fracturing have developed into linear ridges and valleys due to surficial processes since the beginning of the Adirondack Uplift, approximately 10 - 20 M.a., and continuing to present. The bedrock in this region consists of Precambrian metamorphosed Grenville sedimentary rocks such as Proterozoic marbles, calc-silicates and gneisses that are interlayered with igneous/meta-igneous rocks including granite, gabbro, syenite, anorthosite, gneiss, and diorite (Staaz et al., 1980). The remnants of large tailings piles from historical iron mining operations are scattered throughout this region and used for gravel including the construction of roads. The iron ore bodies are primarily magnetite deposits, including magnetite and apatite, with gangue minerals such as augite, hornblende, albite and tourmaline, and many of these minerals have been shown to host a variety of rare earth elements (Mckeown and Klemic, 1957).

Central to the recent geomorphologic development of the Witherbee quadrangle were the continental ice sheets that intermittently advanced and retreated across the landscape. During the Pleistocene Epoch (2.588 M.a. to 11,700 years B.P.), these cyclic episodes of continental ice sheets periodical overrode this region and were major catalysts in sculpting the present-day landscape by means of their tremendous erosional and depositional nature. Subsequently, the development and presence of proglacial Lake Coveville during the most recent ice sheet retreat inundated the Witherbee quadrangle up to elevations of at least 400 m. For this reason, lacustrine sediment deposits including deltaic structures and shoreline terraces can be observed throughout the eastern portion of this quadrangle.

Dominant surficial materials throughout this mapping area were diamictons that blanketed the topography with varying thicknesses. The diamictons studied can be classified as glacial till, colluvium, or stratified lacustrine or deltaic mixed sands and gravels. In several locations, well-sorted sands were observed at excavated gravel pits and construction sites that displayed excellent stratification including a variety of deposition structures such as cross bedding, soft-sediment faulting and folding, grading, asymmetrical ripple marks and imbrication. The majority of stratified sand consisted of alternating layers of predominately dark (garnet, amphibole) and predominantly light colored (quartz) minerals, approximately 0.1 - 0.3 m thick. At some outcrops, alternating beds of sand and gravel approximately 0.5 – 1.0 m thick were observed dipping east. Rarely observed at any outcrop were deposits of well-sorted stratified silts and clays, although fine to very-fine silty/clayey sands were found.

References Cited

Mckeown, F. A. and Klemic, H., 1957, Rare-earth-bearing apatite, at Mineville, Essex County, New York: Bull. U.S. Geol. Sur., 1046B, 9-23.

Staaz, M. H., Hall, R. B., Macke, D. L., Armbrustmacher, T. J., and Brownfield, I. K., 1980, Thorium resources of selected regions in the United States: U.S. Geological Survey Circular 824, 32p.

Cross-Section A-A' Discussion

The geology displayed in cross section A – A' is based on a combination of surficial mapping and subsurface data, and the topographic profile was extracted from a digital elevation model (DEM). The cross section begins approximately 1.5 km north of the northern-most apex of Tracy Road, and trends towards the southeast corner of the quadrangle, ending just beyond Sherman Lake Road. The surficial geology is comprised primarily of Pleistocene diamicton with localized occurrences of surface sands and subsurface sands, silts and clays. Surficial cover is relatively thin throughout the cross section, with a maximum thickness of approximately 20 m to bedrock. Surface topography is heavily dependent on bedrock control.



QUADRANGLE LOCATION



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UTM GRID AND 2016 MAGNETIC NORTH

DECLINATION AT CENTER OF SHEET

prepared by

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Peat, muck, marl, silt, clay or sand deposited in association with wetland environments. Various sediments can be present at transitional boundaries from one facies to another

Pleistocene

Hw



Stratified Sand (Ps)





environments.





Pre-Pleistocene



Bedrock (Br) Non-glacially derived, hard rock, pre-pleistocene in age. May be covered up to a meter in diamicton, sand and gravel, or sand and clay in areas marked as Br.



ADJOINING QUADRANGLES

Poolsy Peak Ridde	Elizabethown	Westport
Underwood	Wittendee	PortHenry
ParadotLake	trapelate	Crown Point



